ATM 265, Spring 2019 Lecture 8 Variable Resolution Modeling April 24, 2019

Paul A. Ullrich (HH 251) paullrich@ucdavis.edu

Slides are based on Colin Zarzycki's talk on variable resolution from the DCMIP workshop (2016)

High-Resolution Modeling

Why do we want higher resolution?

- Improved resolution of land-surface processes (snowpack, runoff)
- Resolution of transient eddies (synoptic-scale frontal systems, local convective systems)
- Resolution of extreme weather events
- Improvement in representation of geographic features (mountain ranges and islands)



The California coastal ranges have a profound effect on regional climate which is poorly captured in current climate models.

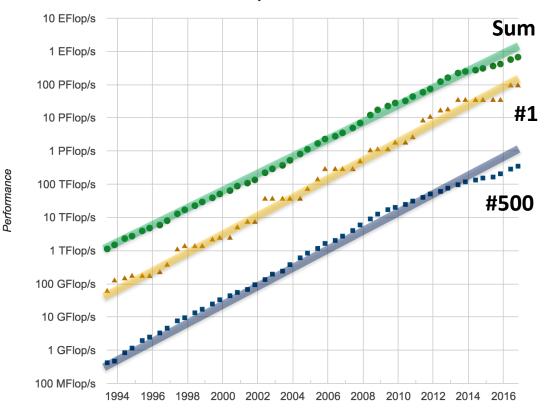
Computing Power vs. Resolution

Computational power doubles approximately every 1.2 years.

To obtain a factor of 2 horizontal refinement, numerical models require 8x the computational power.

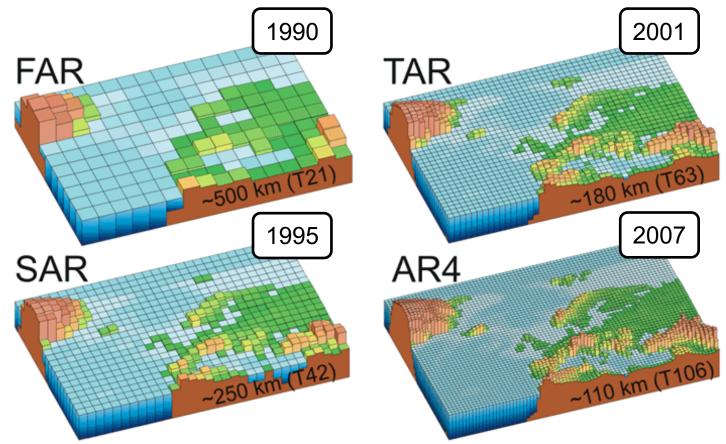


Doubling of horizontal resolution every 3.6 years?



Performance Development

Climate Model Resolution



AR5 (2013) included some model simulations at ~50km, but most runs were at 110km.

AR6 (2019) will rely on CMIP6 runs, which include many runs at 25-50km global resolution.

Computing Power vs. Resolution

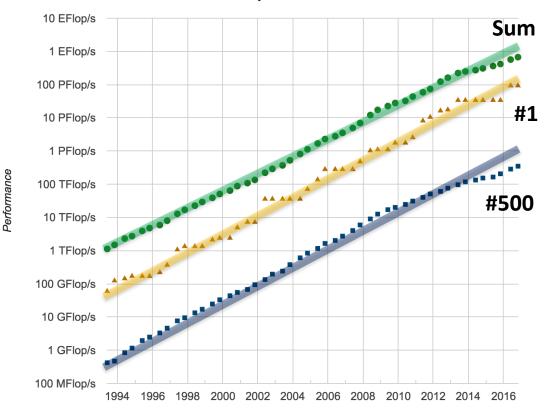


Doubling of horizontal resolution every 3.6 years?

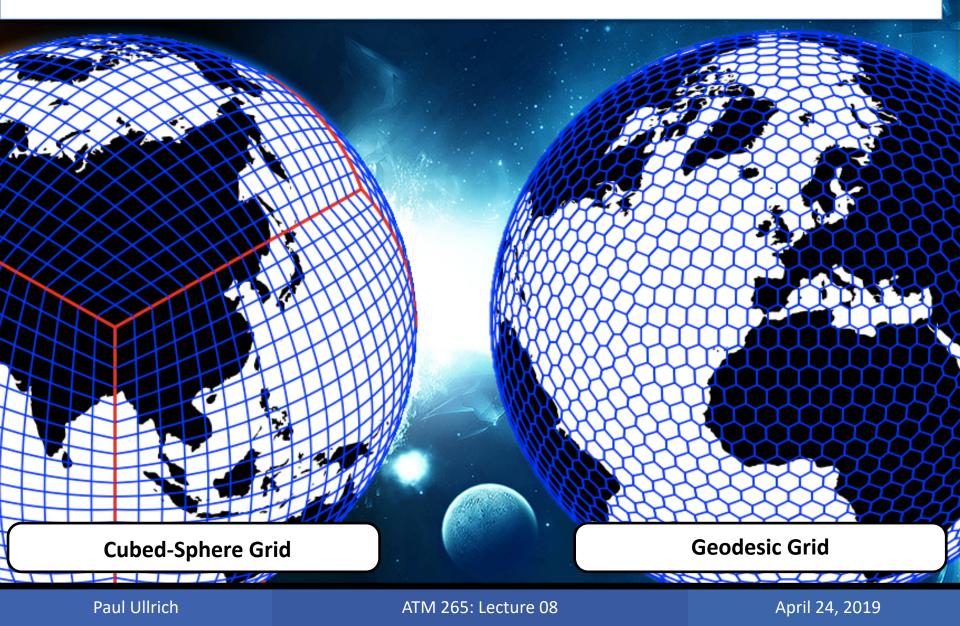
Reality is quite different – doubling of resolution is closer to every 10 years.

Why?

High resolution models are expensive to run. This implies they are hard to tune, and so it is difficult to demonstrate significant improvements in simulation quality. **Performance Development**

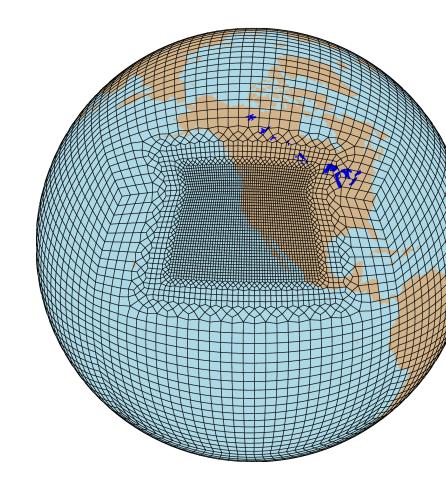


Quasi-Uniform Grids

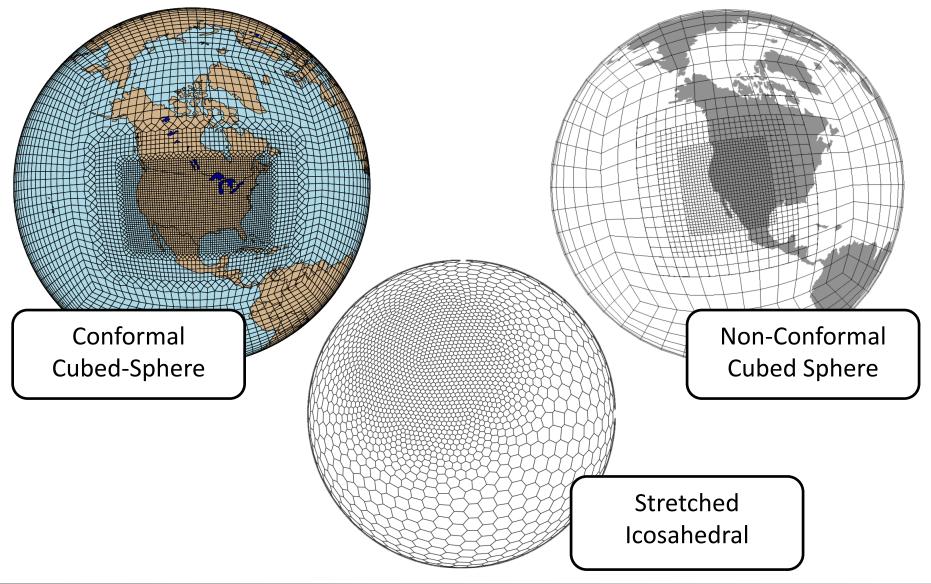


Variable Resolution (VR) Models

- VR allows for **fewer computational resources** to be spent sparingly on a single problem.
- Fully coupled global modeling system, useful for seasonal to subseasonal forecasting.
- More ensemble members can be produced for a particular region (uncertainty quantification).
- Resolution where you need it.



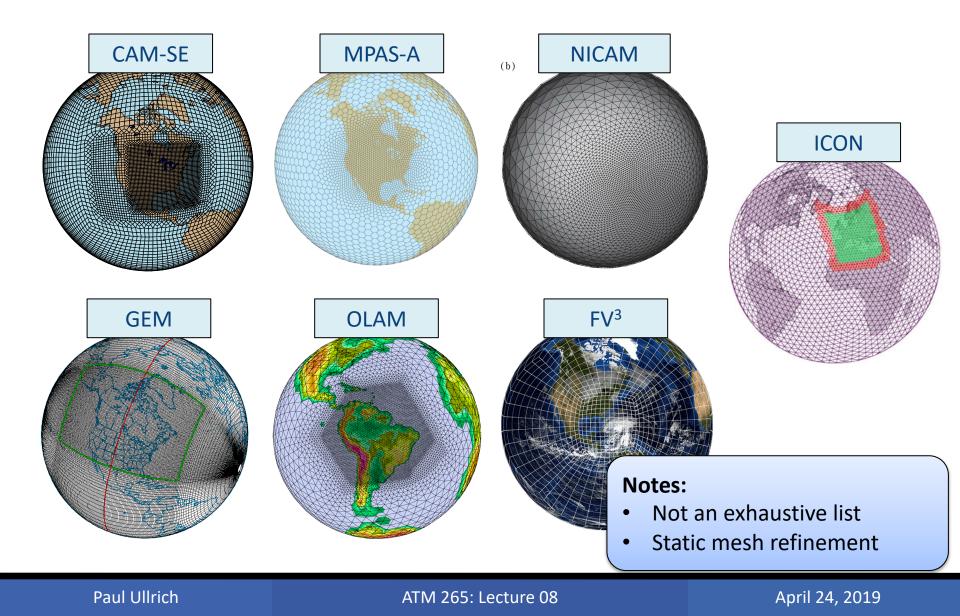
Variable Resolution Models



Paul Ullrich

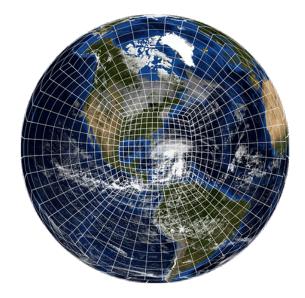
ATM 265: Lecture 08

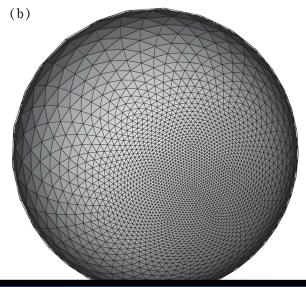
A Sampling of VR Models



Stretched Grids

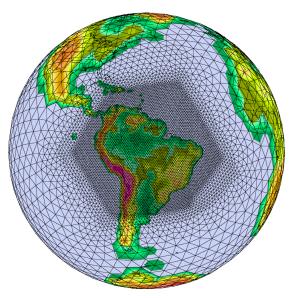
- Examples going back nearly 40 years (e.g., Schmidt, 1977; Staniforth and Mitchell, 1978)
- Generally pole-symmetric dilation
- Benefits
 - Numerical modifications trivial
 - Grid structurally unchanged
- Drawbacks
 - At high stretching factors, far field quickly under-resolved
 - Stretching beyond ~7x problematic (Caian and Geleyn, 1997)

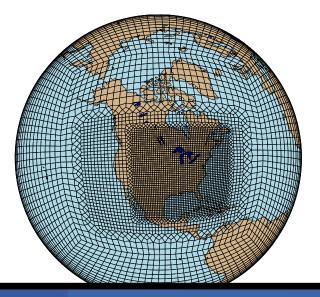




Unstructured / Nested Refinement

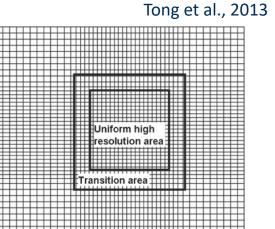
- Many "next-generation" VRGCMs adopting more flexible avenues
 - Add cells in area of interest
- Require more "local" stencils capable of operating on arbitrary grids
- *h*-refinement
- Benefits
 - Doesn't coarsen far-field
 - Multiple regions, flexibility in shape of refinement patches
- Drawbacks
 - Adds cells (cost containment), requires unstructured grids
 - Load balancing, communication (connectivity) may not be as trivial

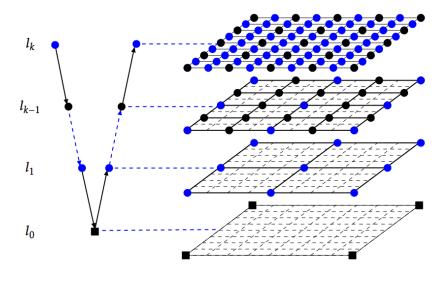




One Other Distinction

- "Single-grid" or "uni-grid" variableresolution
 - Every lat/lon point is covered by one and only one grid cell
 - No remapping/interpolation between grid scales
 - Trivial conservation
- "Multi-grid" variable-resolution
 - More analogous to two-way nesting
 - High-resolution nest "overlays" coarser grid
 - Difference from embedded RGCM? Same model, "single direction" timestepping



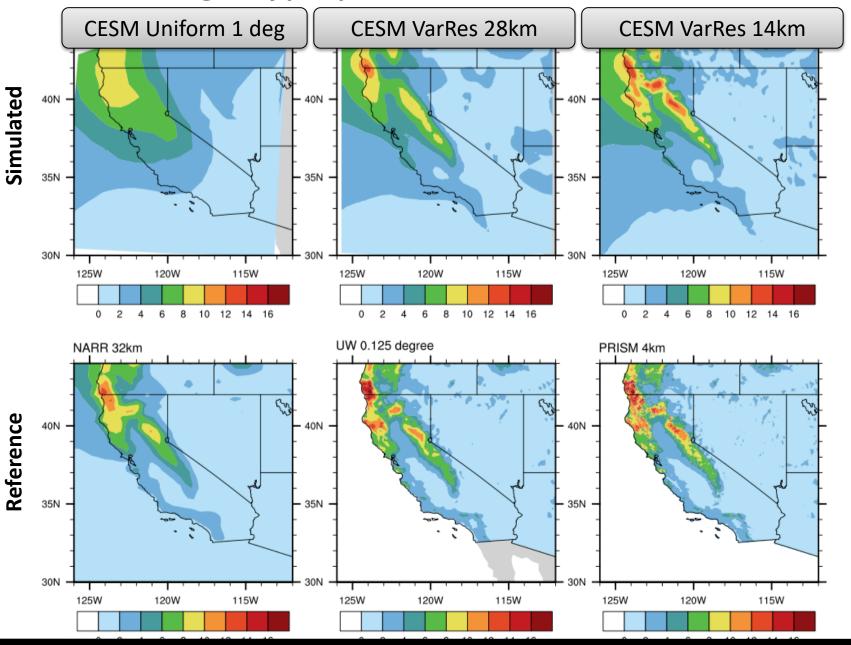


Wenqiang Feng, U. Tenn.

Paul Ullrich

ATM 265: Lecture 08

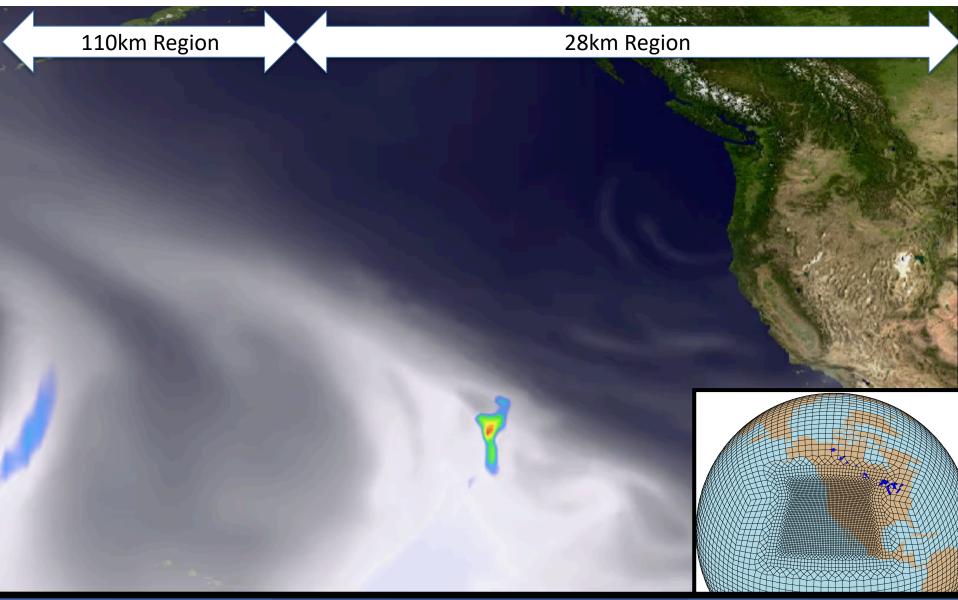
Average daily precipitation rate DJF 1980-1986 unit:mm/d



Paul Ullrich

ATM 265: Lecture 08

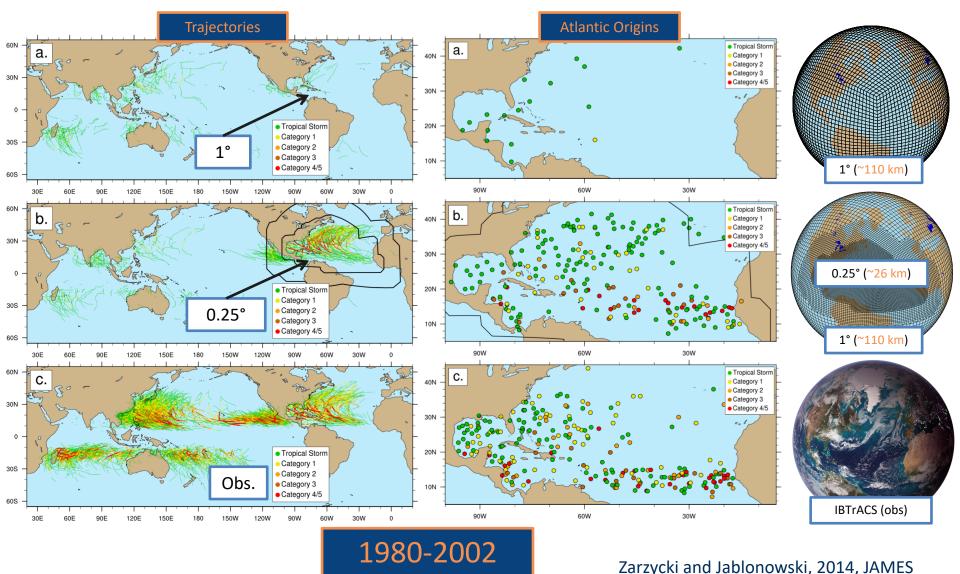
Applications: Atmospheric Rivers



Paul Ullrich

ATM 265: Lecture 08

Applications: Tropical Cyclone Climatology



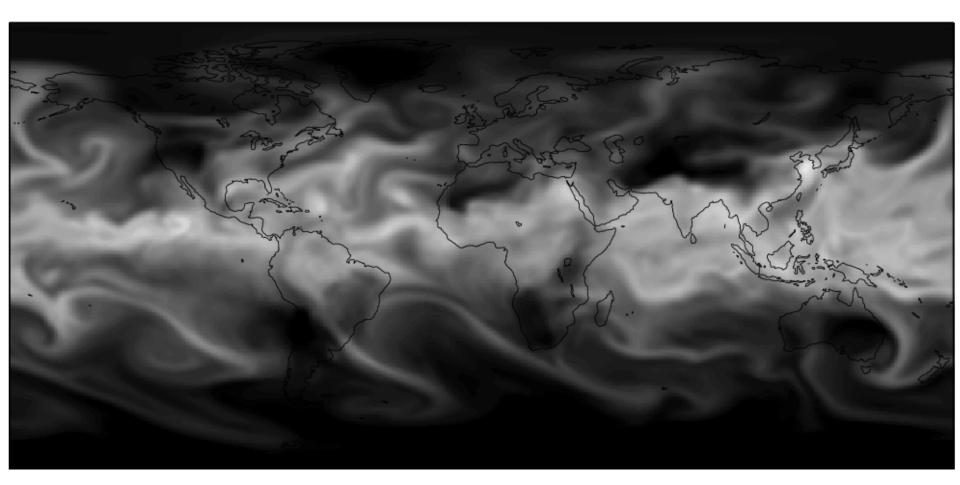
ATM 265: Lecture 08

April 24, 2019

Paul Ullrich

Applications: Tropical Cyclone Climatology

Uniform-Resolution Global Simulation

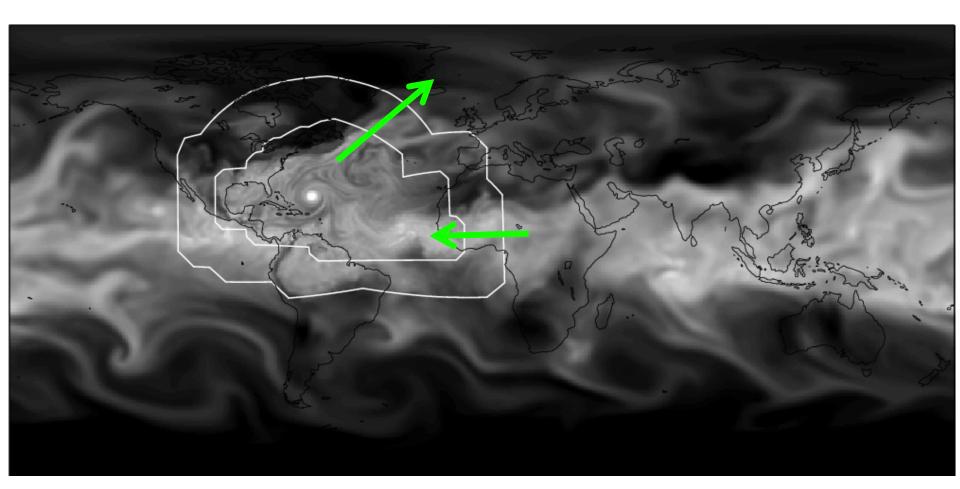


Paul Ullrich

ATM 265: Lecture 08

Applications: Tropical Cyclone Climatology

Uniform-Resolution Global Simulation

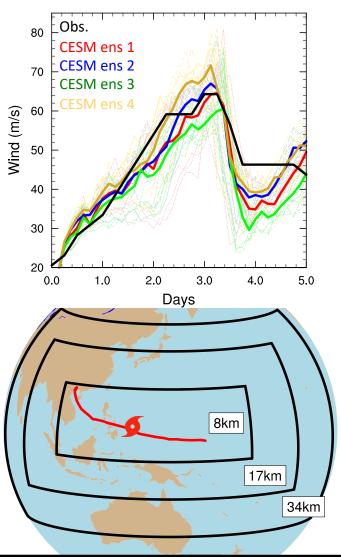


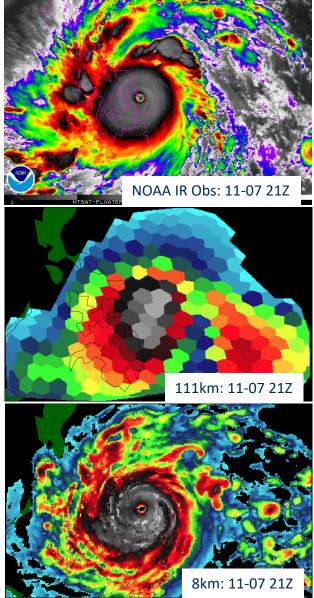
Paul Ullrich

ATM 265: Lecture 08

Applications: Typhoon Haiyan Forecasts

- VR-CESM produces realistic track, intensity and structure
- Computationally inexpensive within coupled earth system model



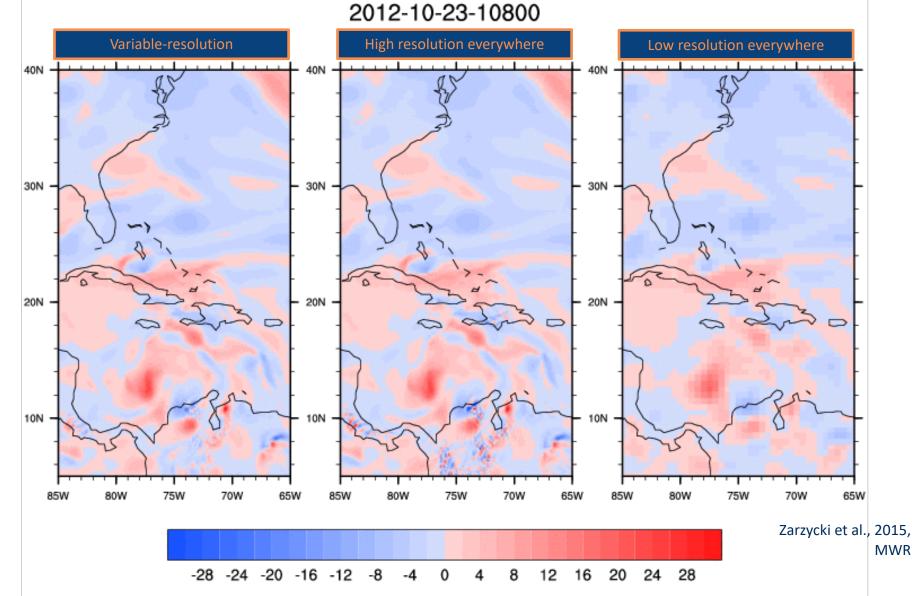


Paul Ullrich

ATM 265: Lecture 08

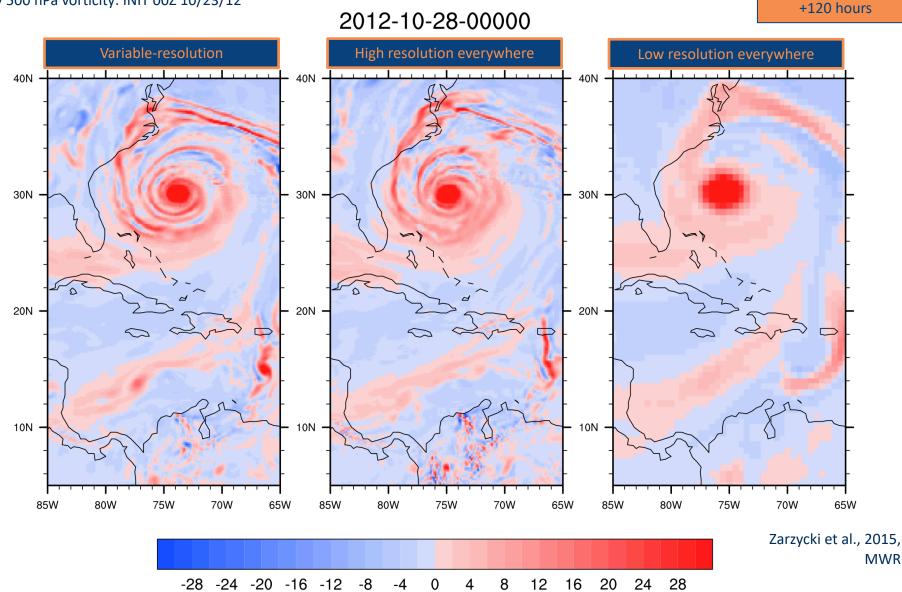
Applications: Hurricane Sandy Forecast

Sandy 500 hPa vorticity: INIT 00Z 10/23/12



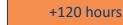
Applications: Hurricane Sandy Forecast

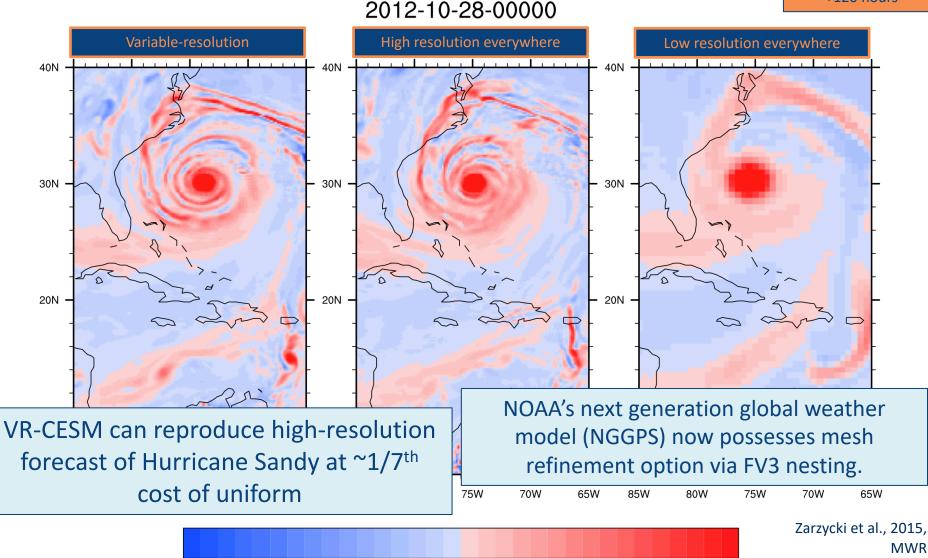
Sandy 500 hPa vorticity: INIT 00Z 10/23/12



Applications: Hurricane Sandy

Sandy 500 hPa vorticity: INIT 00Z 10/23/12





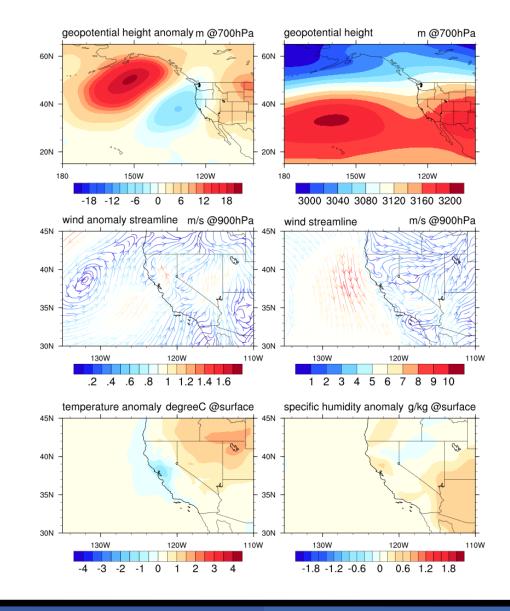
-28 -24 -20 -16 -12 -8 -4 0 4 8 12 16 20 24 28

Applications: Marine Air Intrusion

Central Valley Delta Breeze events are important for cooling and ventilating the central valley, and bringing relief from heat waves

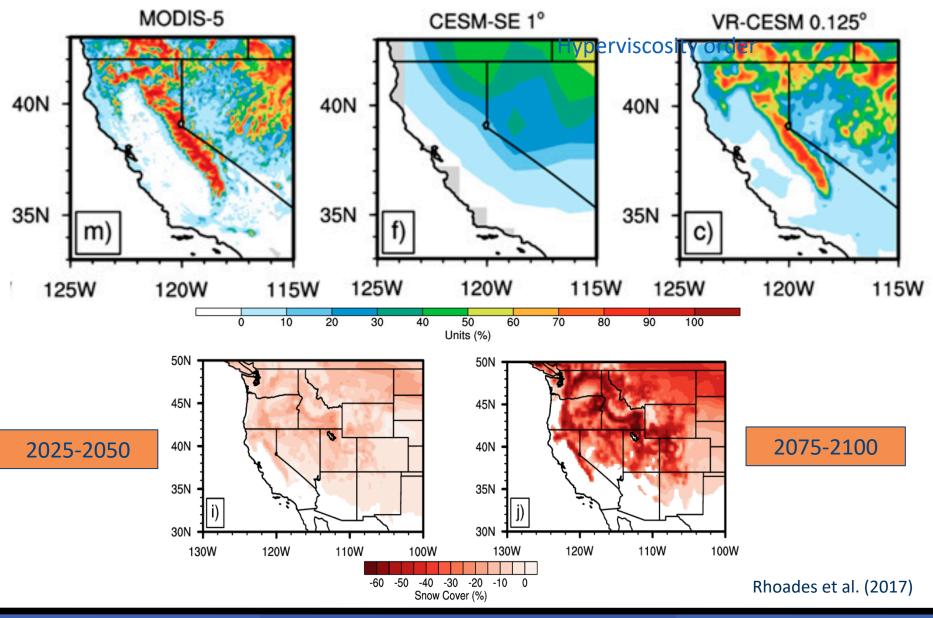
These are regional meteorological features driven by large-scale meteorological patterns (LSMPs).

Variable-resolution ensembles have been used to isolate the LSMPs associated with Delta Breeze events, allowing them to be predicted based on simulations that only resolve the large-scale flow.



ATM 265: Lecture 08

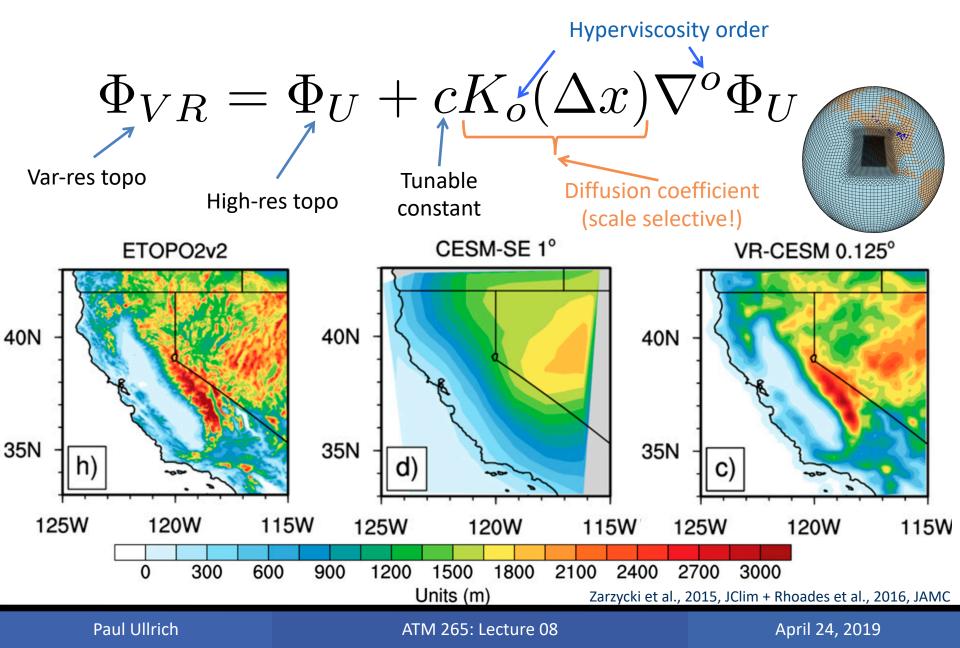
Applications: Snowpack

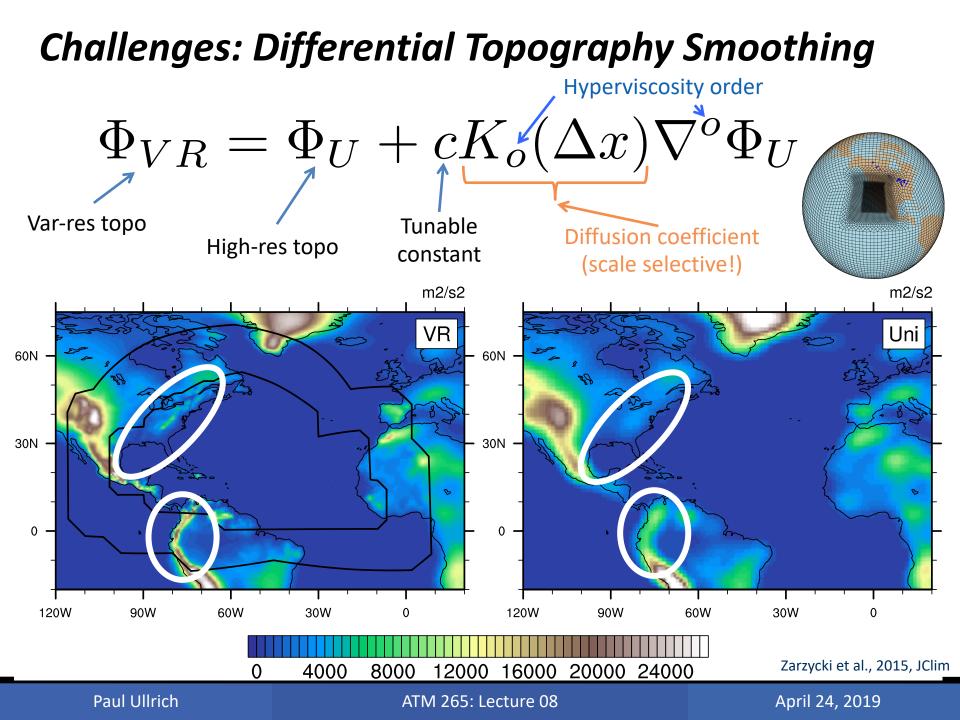


Paul Ullrich

ATM 265: Lecture 08

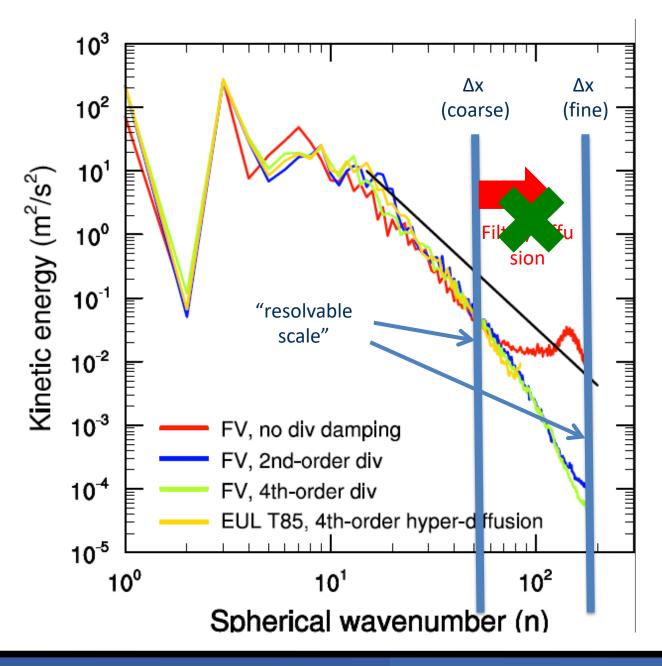
Challenges: Differential Topography Smoothing





Challenges: Diffusion

Has been covered that models generally need some form of diffusion to **remain stable** and **produce realistic results**



Challenges: Diffusion

- Some schemes have enough implicit diffusion that they don't require additional filtering
 - Implicit diffusion inherently scale-selective
- Other models require explicit diffusion for numerical stability and to remove grid-scale noise
 - Smagorinsky
 - Controlled by deformation/stability of local flow
 - (Hyper)-viscosity
 - Applied as forcing term in relevant state equations
- Explicit diffusion requires careful care to <u>only operate on spurious energy</u> <u>near grid scale (e.g., numerical noise, wave reflection, etc.)</u>

Challenges: Diffusion 701

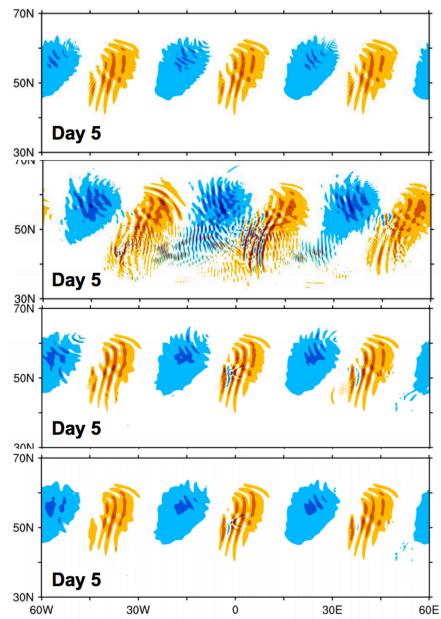
Uniform mesh ($\Delta x = 30$ km) Smagorinsky

TR10 mesh ($\Delta x = 90-30$ km) Smagorinsky, Δx^2 scaling

TR10 mesh ($\Delta x = 90-30$ km) background K₄ = 1x10¹² m⁴s⁻¹ (30 km mesh value, Δx^4 scaling)

TR10 mesh ($\Delta x = 90-30$ km) background K₄ = $3x10^{12}$ m⁴s⁻¹ (30 km mesh value, Δx^4 scaling)

Skamarock, 2012, PDES



Paul Ullrich

ATM 265: Lecture 08

Challenges: Diffusion

CAM-SE, Zarzycki et al., 2014, JClim

Paul Ullrich

$$K_{4}(\Delta x) = K_{4}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{4}(\Delta x) = K_{4}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{4}(\Delta x) = K_{4}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{4}(\Delta x) = K_{4}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{5}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{6}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

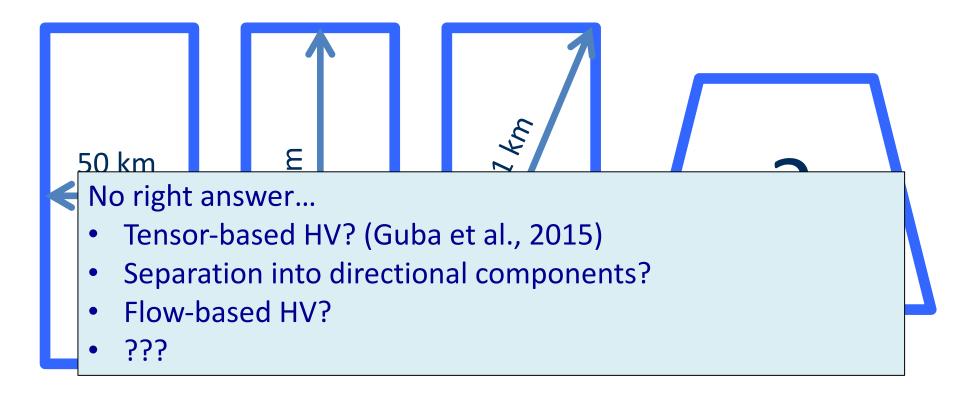
$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

$$K_{7}(\Delta x_{ref}) \begin{pmatrix} \Delta x \\ \Delta x_{ref} \end{pmatrix}^{y}$$

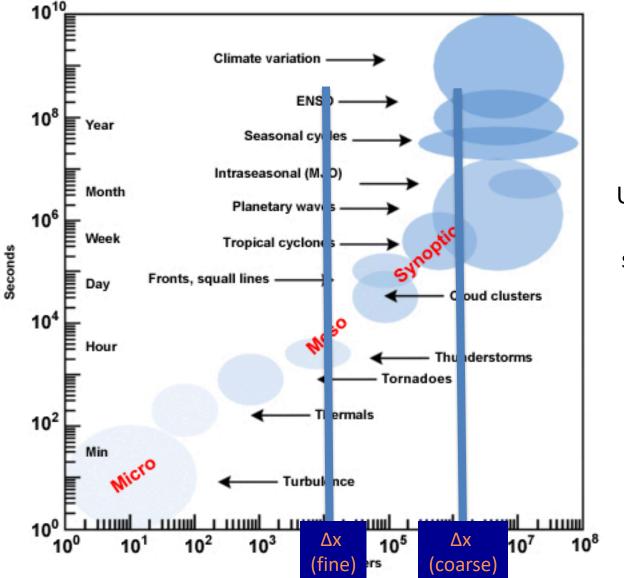
$$K_{7}(\Delta x_{re}$$

Challenges: Diffusion

- Works well for meshes with undistorted elements
- But what happens when you have odd shapes?



Challenges: Scale-Aware Physics

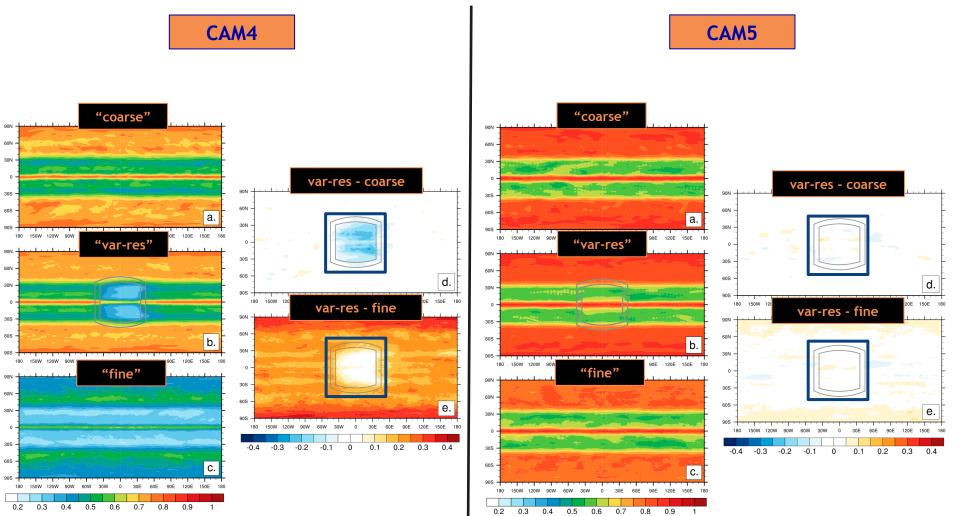


Unified model setups need to be able to properly estimate subgrid properties in all cells

Challenging when spanning large spatial scale gaps...

©The COMET Program

Challenges: Scale-Aware Physics



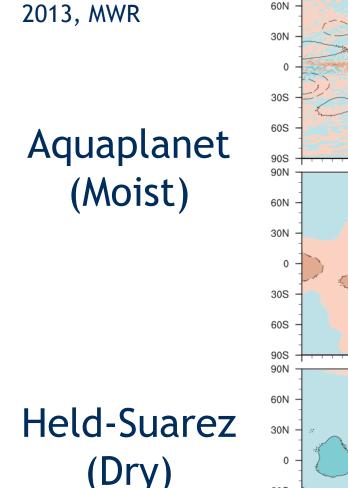
Mean annual cloud fraction

Zarzycki et al., 2014, JClim

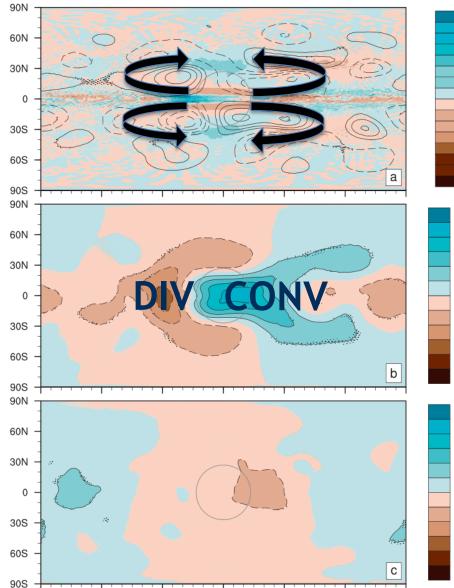
Paul Ullrich

ATM 265: Lecture 08

Challenges: Scale-Aware Physics



Rauscher et al.,



Precip. anom./ Streamlines

6

3

0 -1

-3

-6

0.25

0.2 0.15

0.1

0.05

0 -0.05

-0.1 -0.15 -0.2

-0.25

0.25

0.2

0.15

0.1

0.05

0

-0.05

-0.1

-0.2

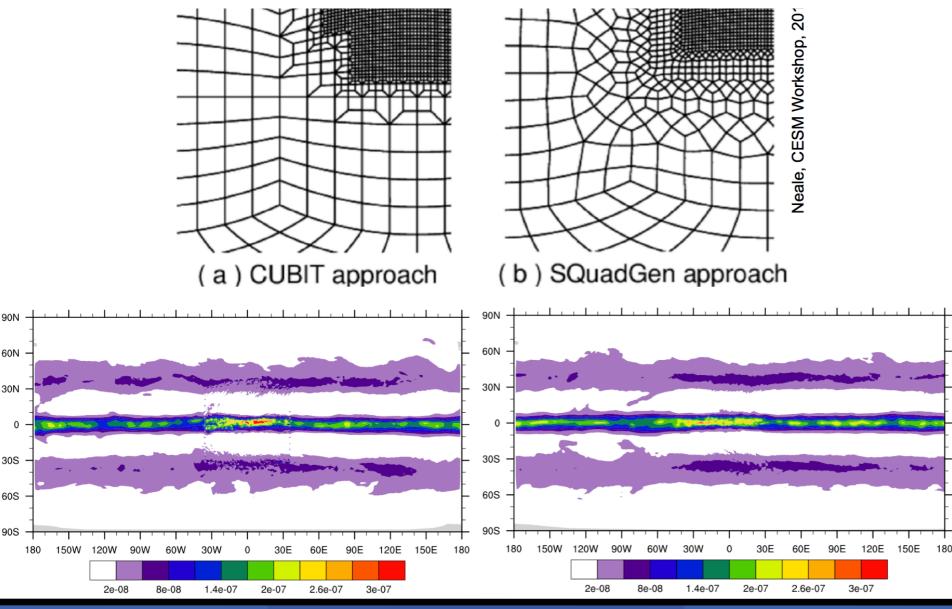
-0.25

200 hPa Eddy Velocity Potential 200 hPa

eddy Velocity Potential

ATM 265: Lecture 08

Challenges: Grid Generation



Paul Ullrich

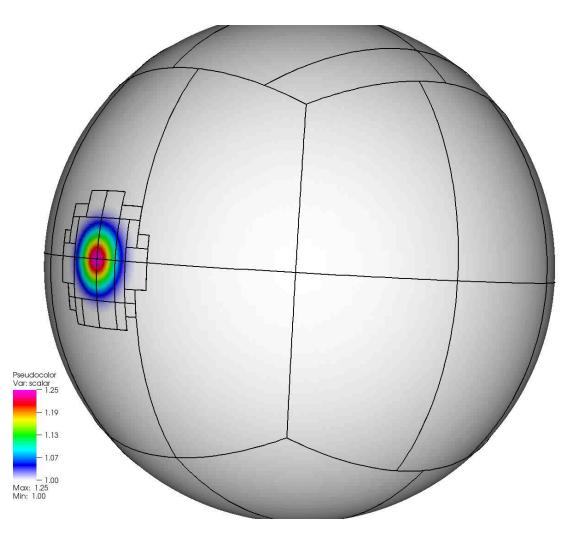
ATM 265: Lecture 08

Adaptive Mesh Refinement (AMR)

With Hans Johansen, Phillip Colella and many others (LBNL)

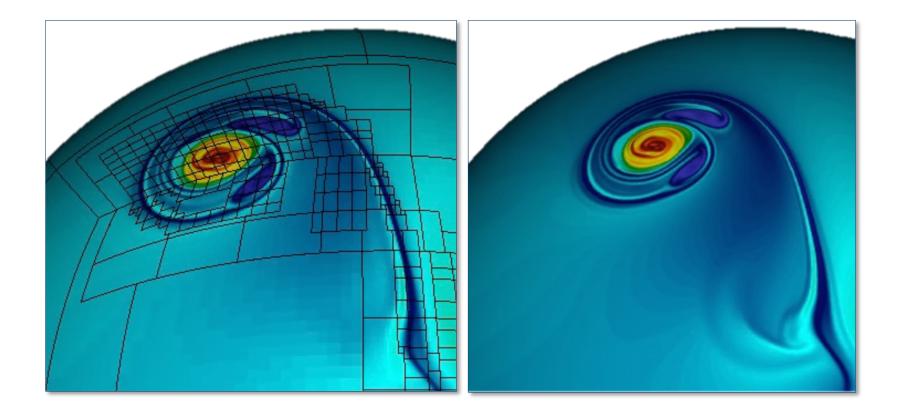
Features of interest can be tracked via adaptive mesh refinement (AMR), which automatically places additional refinement in regions of interest.

In this case, we use the Chombo global circulation model, developed at Lawrence Berkeley National Lab.



Adaptive Mesh Refinement (AMR)

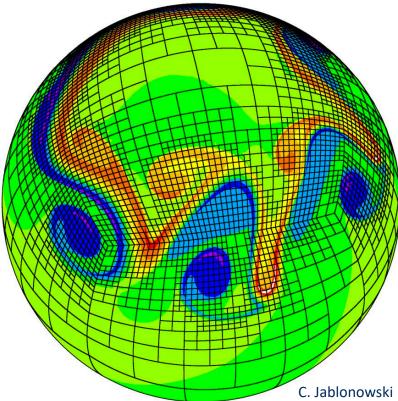
Adaptive mesh refinement poses a particular challenge for future exascale applications due to the need for dynamic load-balancing.



Adaptive Mesh Refinement (AMR)

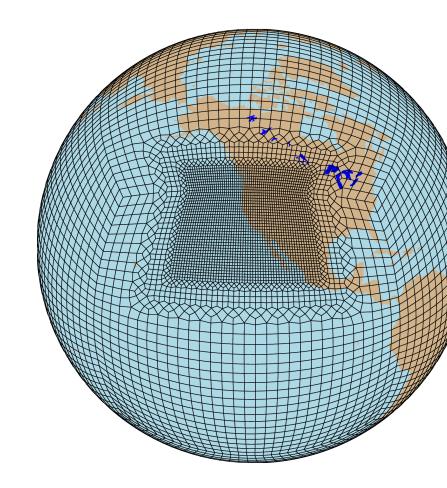
AMR introduces new challenges

- How to "tag" regions to refine (don't want to under-refine or over-refine)
- Load balancing on parallel computing systems
- Need to be able to change configuration "on the fly"
 - Topography
 - Diffusion
 - Sub-grid physics
 - Etc.



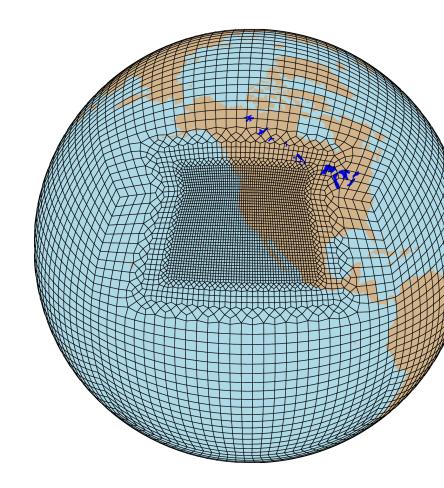
Summary: Variable Resolution Models

- VR allows for **fewer computational resources** to be spent sparingly on a single problem.
- Fully coupled global modeling system, usable for seasonal to subseasonal forecasting.
- More ensemble members can be produced for a particular region (uncertainty quantification).
- Resolution where you need it.



Summary: Variable Resolution Models

- Variable-resolution dynamical cores offer ability to have fine regional resolution in a global modeling framework
- Demonstrated fidelity with tropical cyclones, orographic precipitation, mesoscale convection
- New challenges (hint: avenues for research!)
 - Numerical techniques
 - Scale-aware physics
 - Grid and refinement choices
 - Software engineering



Paul Ullrich