The background of the slide is a vibrant space scene. On the left, a large, dark planet with a textured surface is partially visible. In the center, a bright blue star or nebula glows, with a smaller, blue-tinted planet orbiting it. The right side of the image is filled with a dense field of blue and white stars, creating a sense of depth and cosmic wonder. Two white rounded rectangular boxes with black borders are overlaid on the image, containing text.

Applications of the Basic Equations Chapter 3

Paul A. Ullrich
paulrich@ucdavis.edu

Part 3: The Thermal Wind

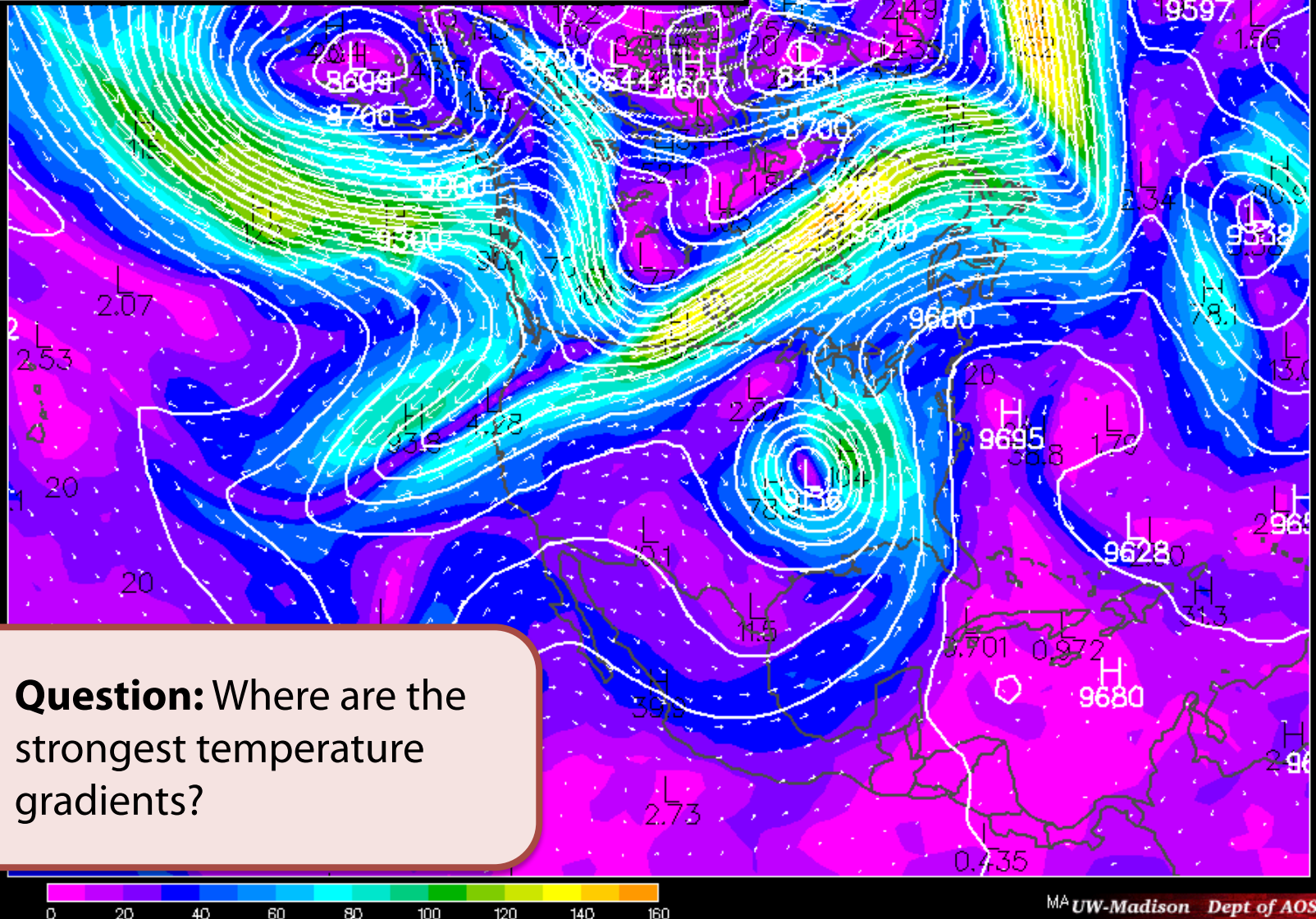


Question: Is there a relationship between wind speed and temperature?

It turns out that there is a close link between **vertical wind shear** (vertical gradients of horizontal wind speed) and **layer thickness**, which is governed by temperature.

300 hPa Wind Speed

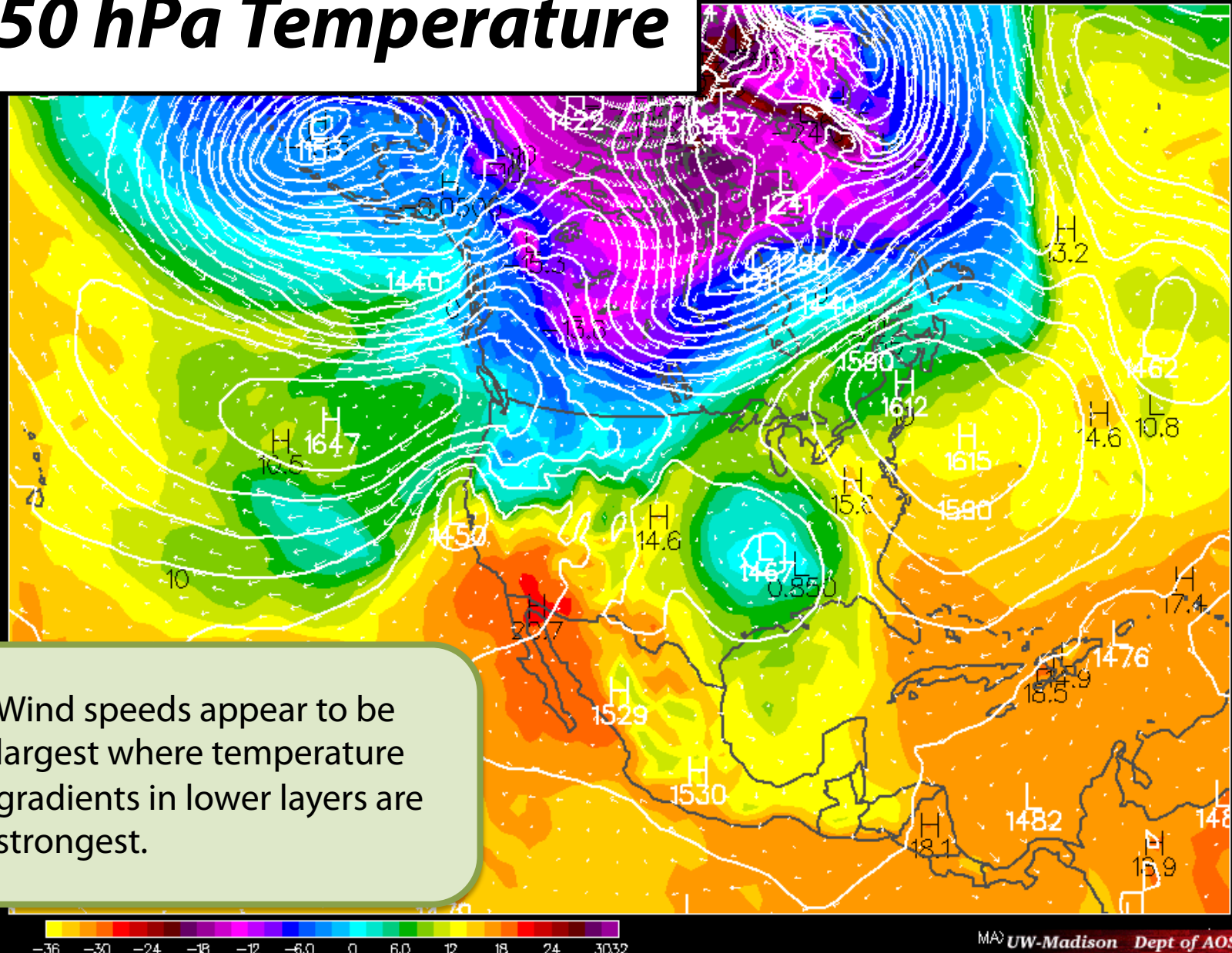
12 hour GFS valid 12Z FRI 26 OCT 07



Question: Where are the strongest temperature gradients?

850 hPa Temperature

12 hour GFS valid 12Z FRI 26 OCT 07



Wind speeds appear to be largest where temperature gradients in lower layers are strongest.

Thermal Wind

Definition: The **thermal wind** is a vector difference between the geostrophic wind at an upper level and a lower level.

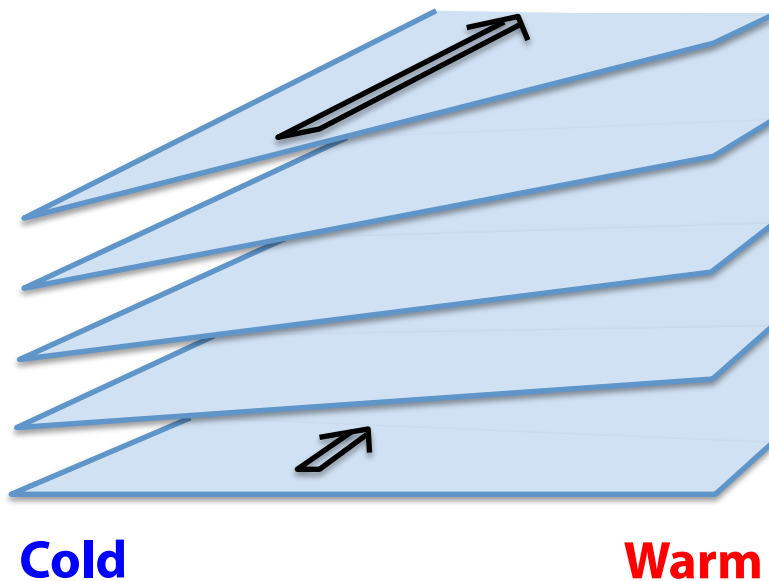
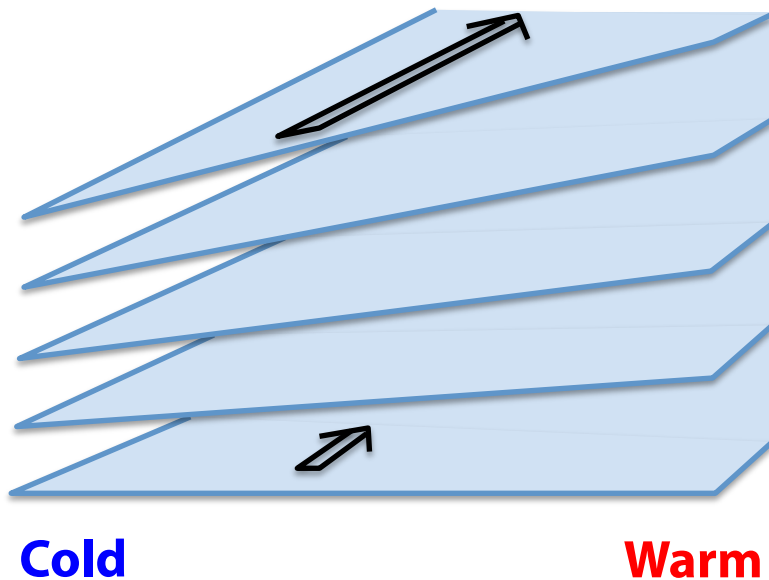


Figure: Thickness of layers related to temperature, causes a tilt in the pressure surfaces.

Change in magnitude of horizontal gradient of pressure then leads to vertical wind shear.

Thermal Wind

Emphasis: The thermal wind is not a real wind, but a vector difference.



The thermal wind vector points such that **cold air is to the left** and **warm air is to the right**, parallel to isotherms (in the northern hemisphere). Cold air is to the right and warm air is to the left in the southern hemisphere.

Geostrophic Wind

In Pressure Coordinates

Recall: The **geostrophic wind** is the component of the real wind which is governed by geostrophic balance. On constant height surfaces, it is defined to satisfy

$$u_g = -\frac{1}{f\rho} \frac{\partial p}{\partial y}$$
$$v_g = +\frac{1}{f\rho} \frac{\partial p}{\partial x}$$

In pressure coordinates, this relationship takes on the analogous form:

$$u_g = -\frac{1}{f} \frac{\partial \Phi}{\partial y}$$
$$v_g = +\frac{1}{f} \frac{\partial \Phi}{\partial x}$$

Geostrophic
Wind

$$u_g = -\frac{1}{f} \left(\frac{\partial \Phi}{\partial y} \right)_p \quad v_g = \frac{1}{f} \left(\frac{\partial \Phi}{\partial x} \right)_p$$

Differentiate with respect to p



$$\frac{\partial u_g}{\partial p} = -\frac{1}{f} \frac{\partial}{\partial y} \frac{\partial \Phi}{\partial p} \quad \frac{\partial v_g}{\partial p} = \frac{1}{f} \frac{\partial}{\partial x} \frac{\partial \Phi}{\partial p}$$

Hydrostatic
Relationship

$$\frac{\partial \Phi}{\partial p} = g \frac{\partial z}{\partial p} \downarrow = -\frac{1}{\rho} = -\frac{RT}{p}$$

On constant p
surfaces

**Thermal wind
relationship:**

$$\frac{\partial u_g}{\partial p} = \frac{R}{pf} \left(\frac{\partial T}{\partial y} \right)_p \quad \frac{\partial v_g}{\partial p} = -\frac{R}{pf} \left(\frac{\partial T}{\partial x} \right)_p$$

Thermal wind relationship: $\frac{\partial u_g}{\partial p} = \frac{R}{pf} \left(\frac{\partial T}{\partial y} \right)_p$ $\frac{\partial v_g}{\partial p} = -\frac{R}{pf} \left(\frac{\partial T}{\partial x} \right)_p$


This relationship links the horizontal temperature gradient with the vertical wind gradient (vertical shear).


Thermal Wind

Thermal wind relationship: $\frac{\partial u_g}{\partial p} = \frac{R}{pf} \left(\frac{\partial T}{\partial y} \right)_p$

$$\frac{\partial v_g}{\partial p} = -\frac{R}{pf} \left(\frac{\partial T}{\partial x} \right)_p$$

The thermal wind itself is a vector difference.

Rewrite  $\frac{\partial u_g}{\partial(\log p)} = \frac{R}{f} \left(\frac{\partial T}{\partial y} \right)_p$ $\frac{\partial v_g}{\partial(\log p)} = -\frac{R}{f} \left(\frac{\partial T}{\partial x} \right)_p$

Integrate  $u_T = u_g(p_2) - u_g(p_1) = -\frac{R}{f} \left(\frac{\partial \langle T \rangle}{\partial y} \right)_p \log \left(\frac{p_1}{p_2} \right)$

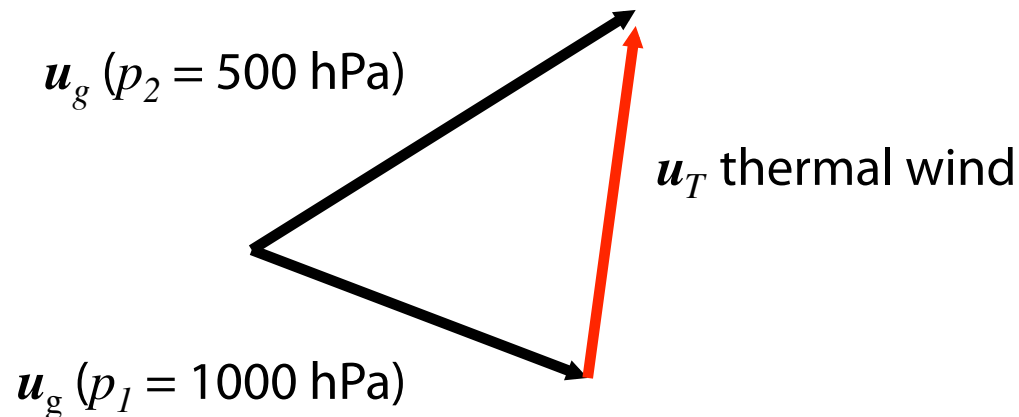
Thermal Wind $v_T = v_g(p_2) - v_g(p_1) = \frac{R}{f} \left(\frac{\partial \langle T \rangle}{\partial x} \right)_p \log \left(\frac{p_1}{p_2} \right)$

Thermal Wind

$$u_T = u_g(p_2) - u_g(p_1) = -\frac{R}{f} \left(\frac{\partial \langle T \rangle}{\partial y} \right)_p \log \left(\frac{p_1}{p_2} \right)$$

$$v_T = v_g(p_2) - v_g(p_1) = \frac{R}{f} \left(\frac{\partial \langle T \rangle}{\partial x} \right)_p \log \left(\frac{p_1}{p_2} \right)$$

Example: Thermal wind v_T between 500 hPa and 1000 hPa



Thermal Wind

Alternate form of thermal wind, written in terms of geopotential height (obtained from hypsometric equation):

$$\mathbf{u}_T = \frac{R_d}{f} \mathbf{k} \times \nabla_p \langle T \rangle \log \left(\frac{p_1}{p_2} \right) = \frac{1}{f} \mathbf{k} \times \nabla_p (\Phi_2 - \Phi_1)$$

$$u_T = -\frac{1}{f} \frac{\partial}{\partial y} (\Phi_2 - \Phi_1) \quad v_T = \frac{1}{f} \frac{\partial}{\partial y} (\Phi_2 - \Phi_1)$$

Index "1" indicates lower level, "2" indicates upper level.

Thermal Wind

$$u_T \approx u_g(p_2) - u_g(p_1) = -\frac{R}{f} \left(\frac{\partial \langle T \rangle}{\partial y} \right)_p \log \left(\frac{p_1}{p_2} \right)$$

$$v_T \approx v_g(p_2) - v_g(p_1) = \frac{R}{f} \left(\frac{\partial \langle T \rangle}{\partial x} \right)_p \log \left(\frac{p_1}{p_2} \right)$$

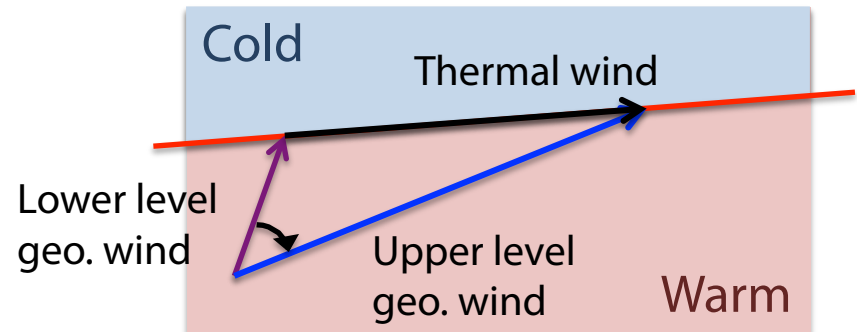
Note that thermal wind always points parallel to lines of constant layer temperature:

$$\mathbf{u}_T \cdot \nabla \langle T \rangle = \frac{R}{f} \log \left(\frac{p_1}{p_2} \right) \left[-\frac{\partial \langle T \rangle}{\partial y} \frac{\partial \langle T \rangle}{\partial x} + \frac{\partial \langle T \rangle}{\partial x} \frac{\partial \langle T \rangle}{\partial y} \right] = 0$$

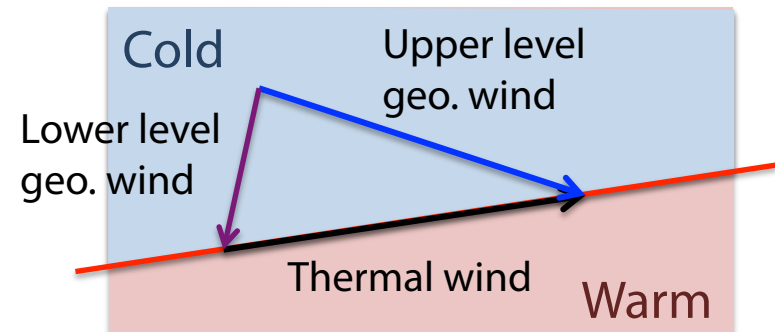
Thermal Wind

Thermal wind always points parallel to lines of constant temperature (and lines of constant layer thickness).

Definition: Veering winds are defined by clockwise rotation of the geostrophic wind with height and are associated with **warm air advection**.



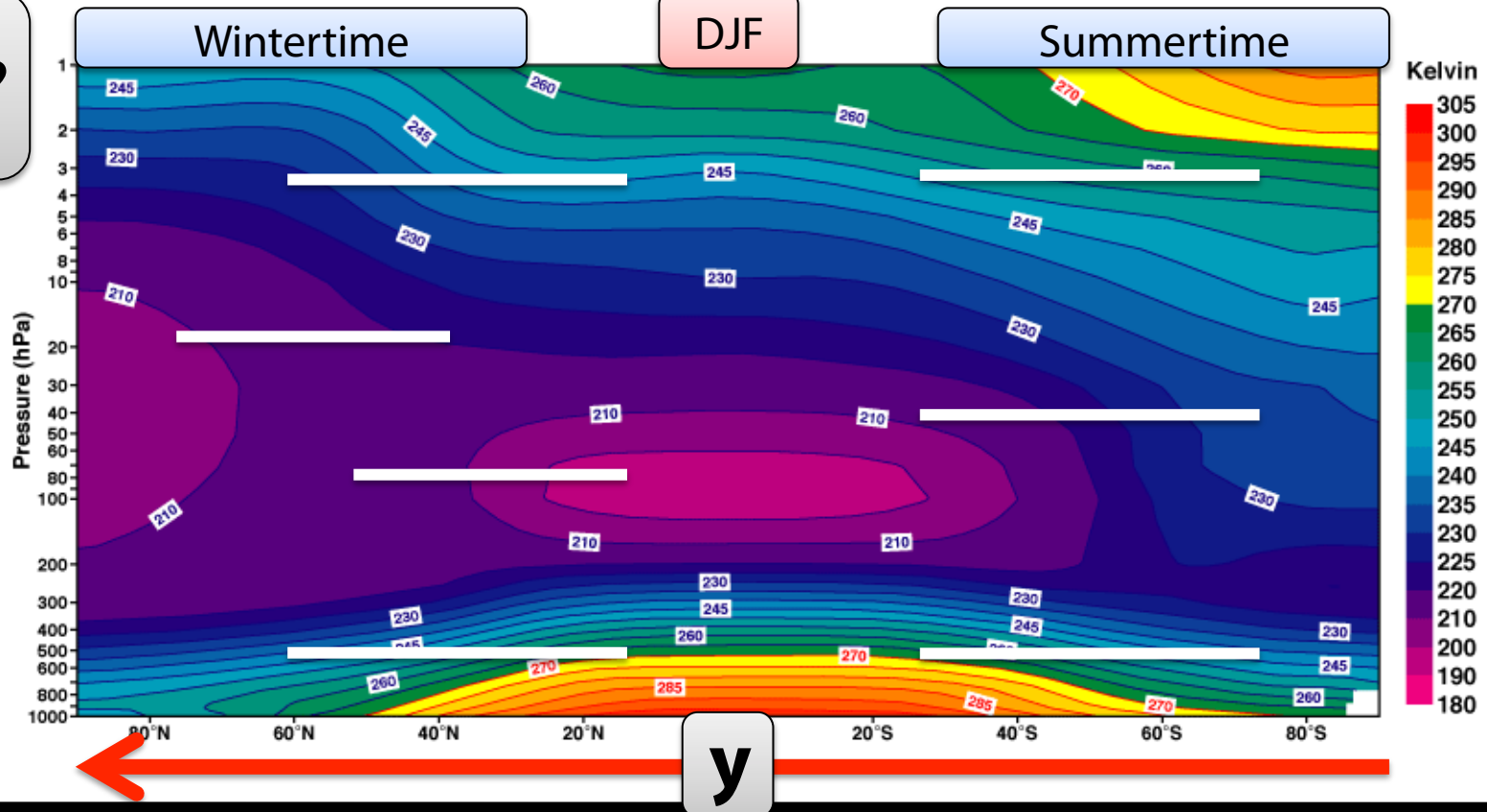
Definition: Backing winds are defined by counter-clockwise rotation of the geostrophic wind with height and are associated with **cold air advection**.



The thermal wind determines the relationship between meridional temperature gradients and zonal winds.

Question: Given zonal mean temperature below, where are zonal jets?

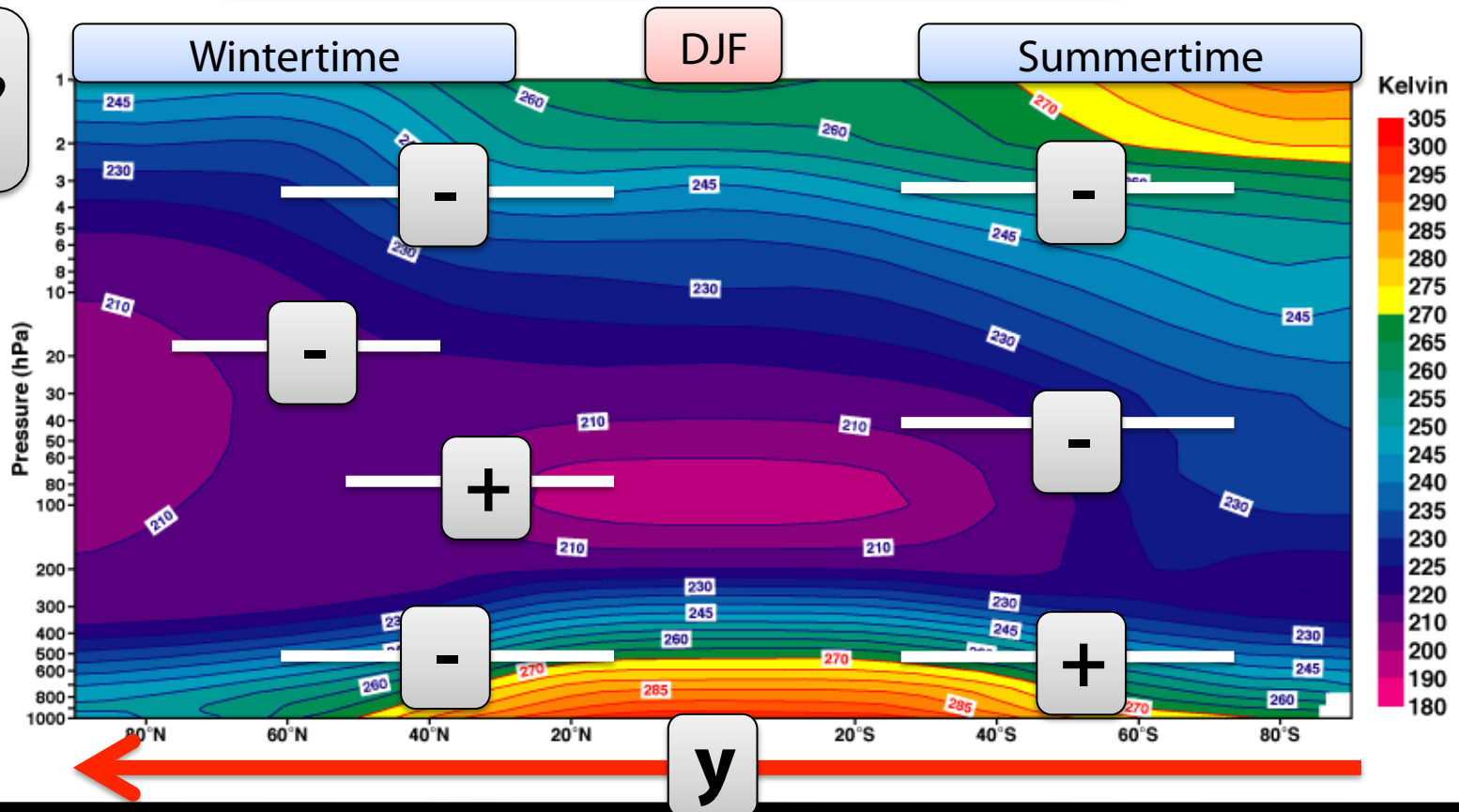
$$\frac{\partial T}{\partial y} ?$$



The thermal wind determines the relationship between meridional temperature gradients and zonal winds.

Question: Given zonal mean temperature below, where are zonal jets?

$$\frac{\partial T}{\partial y} ?$$

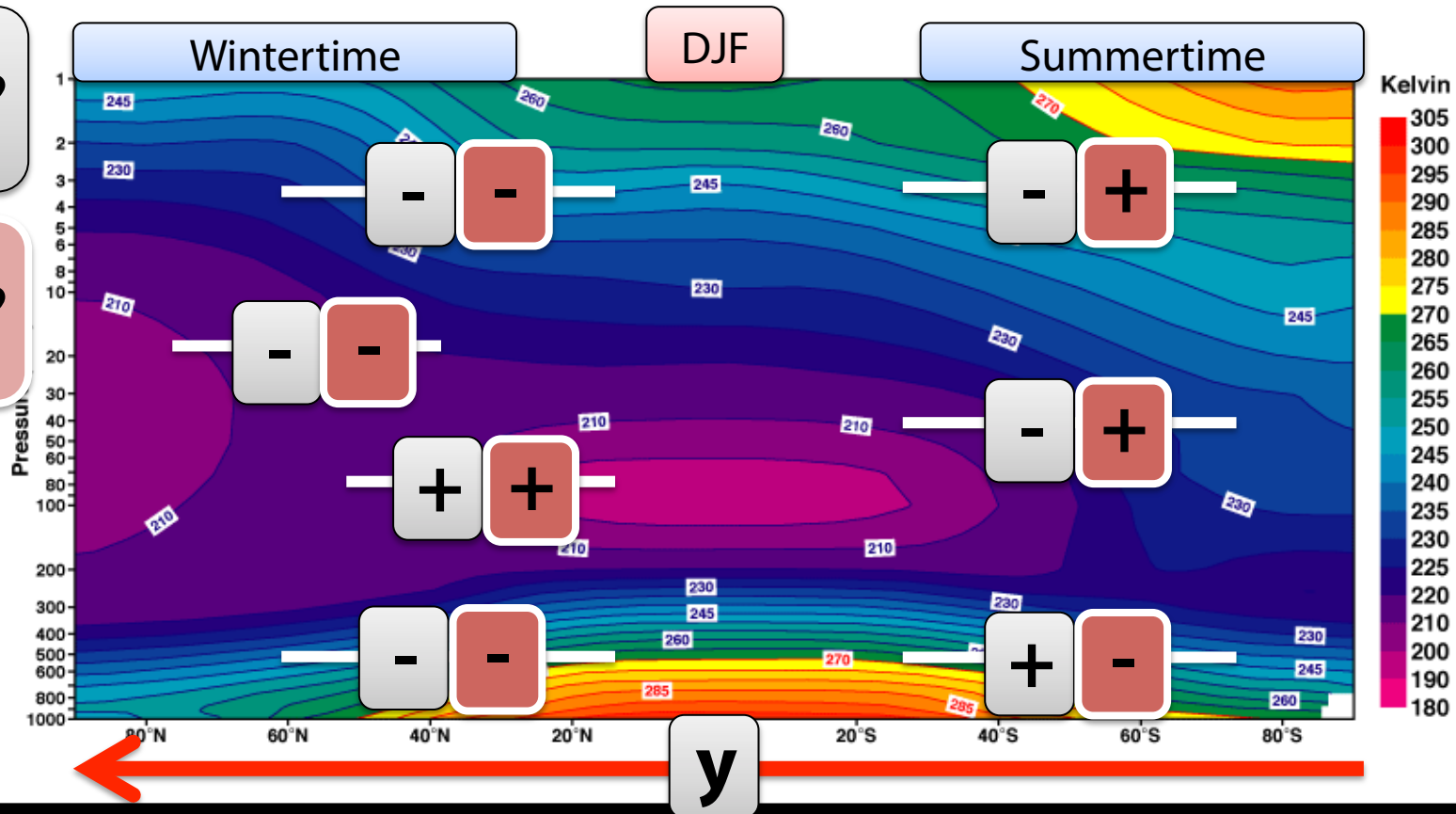


Question: Given zonal mean temperature below, where are zonal jets?

$$\frac{\partial u_g}{\partial p} = \frac{R}{pf} \left(\frac{\partial T}{\partial y} \right)_p$$

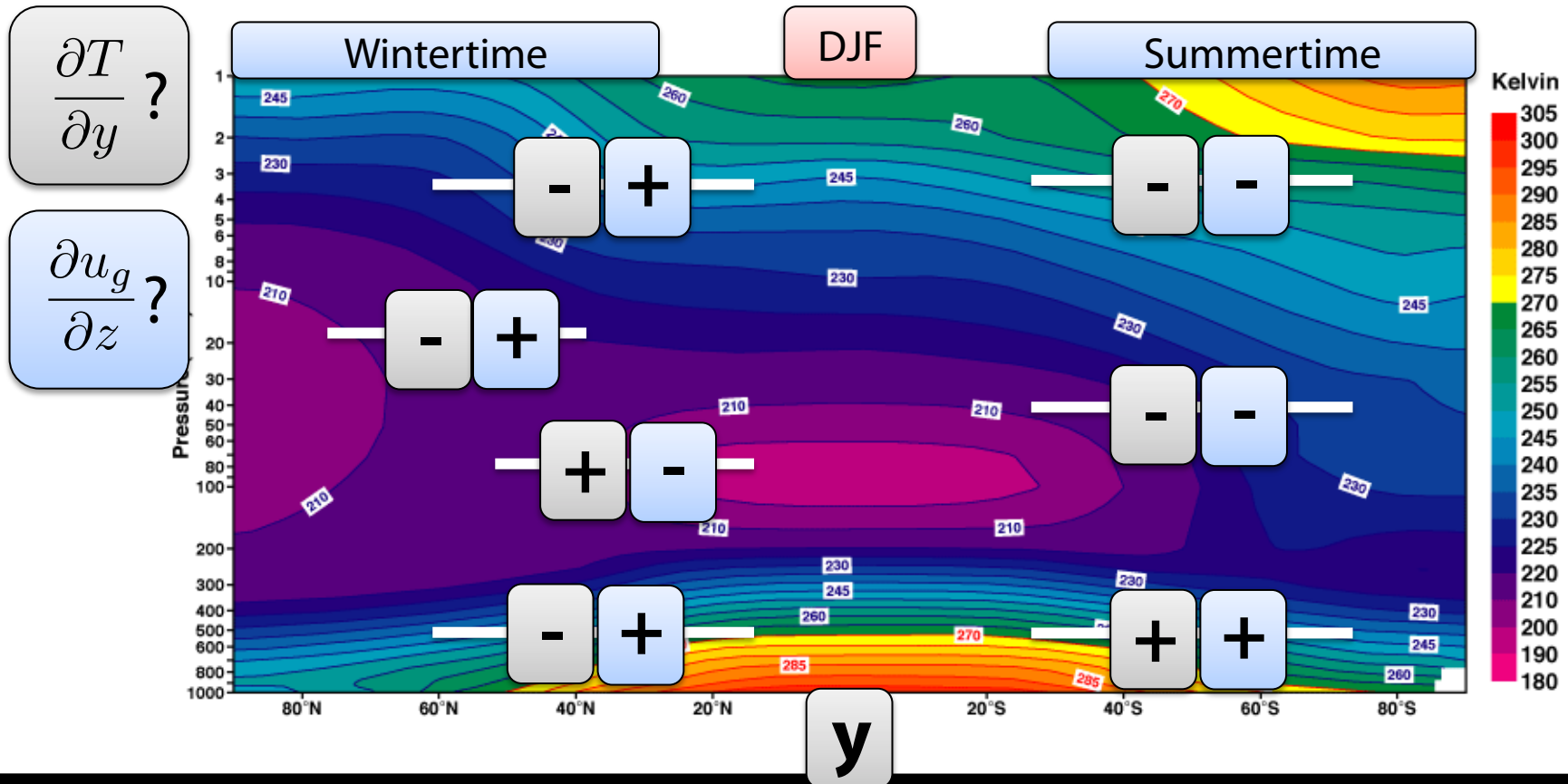
$$\frac{\partial T}{\partial y} ?$$

$$\frac{\partial u_g}{\partial p} ?$$



Question: Given zonal mean temperature below, where are zonal jets?

$$\frac{\partial u_g}{\partial p} = \frac{R}{pf} \left(\frac{\partial T}{\partial y} \right)_p \quad \frac{\partial u_g}{\partial z} \sim - \frac{\partial u_g}{\partial p}$$



Question: Given zonal mean temperature below, where are zonal jets?

$$\frac{\partial u_g}{\partial p} = \frac{R}{pf} \left(\frac{\partial T}{\partial y} \right)_p \quad \frac{\partial u_g}{\partial z} \sim - \frac{\partial u_g}{\partial p}$$

